A 3-D simulation of crustal deformation accompanied by subduction in the Tokai region, central Japan

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Why did we study the Tokai region? Early studies and technical issues

- Modol
- 2. Model

Rate- and state-dependent friction law

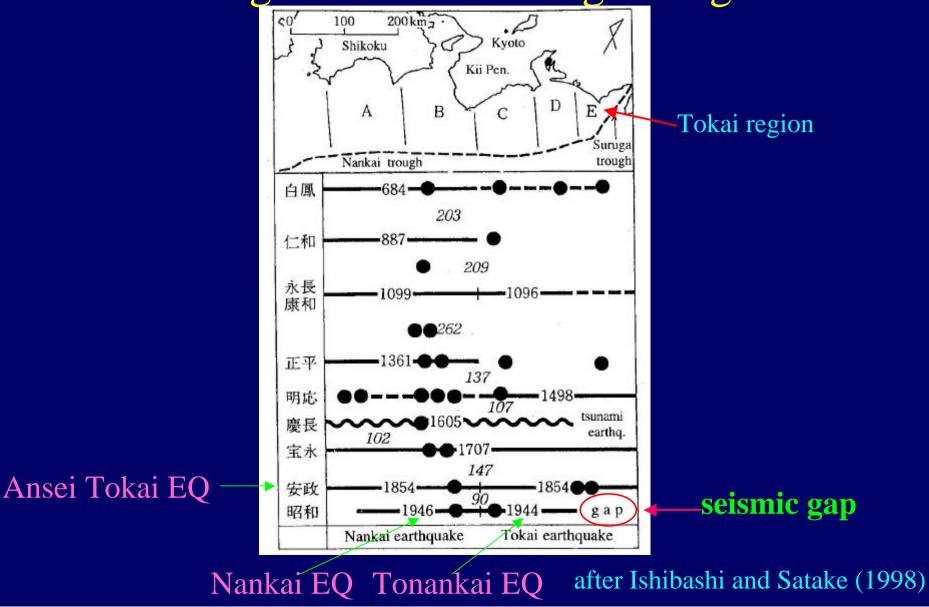
Geophysical constraints by observations

3. Results

Crustal deformation derived from this model

- Comparisons between observation and simulation results
- 4. Summary

Space-time distribution of great earthquakes along the Nankai-Suruga trough



Purpose of our study

Whether precursory changes for the anticipated Tokai earthquake could be observed or not.

Simulation study

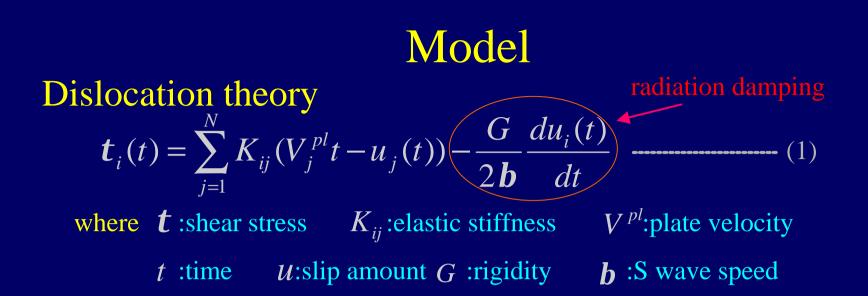
1. Early studies

Kato and Hirasawa (1999) Rate- and state-dependent friction law (Dieterich,1979; Ruina, 1983)

2. Problems

2-D model

3. What's new in the present work3-D model with a curved plate interface



Rate- and state-dependent friction law

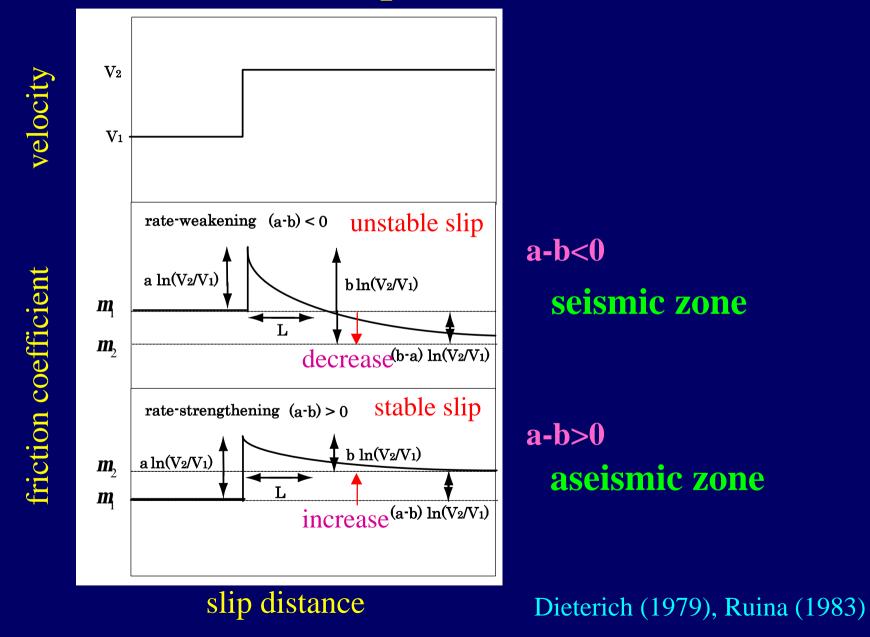
$$\boldsymbol{t}_{i}(t) = \boldsymbol{m}_{i}(t)\boldsymbol{s}_{i}^{eff}$$

$$\boldsymbol{m}_{i}(t) = \boldsymbol{m}_{*} + \boldsymbol{q}_{i}(t) + a_{i}\ln(V_{i}/V_{*})$$

$$\frac{d\boldsymbol{q}_{i}(t)}{dt} = -\frac{V_{i}(t)}{L_{i}}(\boldsymbol{q}_{i}(t) + b_{i}\ln(V_{i}(t)/V_{*}))$$
(2)

where t :frictional forcem:friction coefficientS:normal stressq :state variableV:velocitya, b, L:friction parametersafter Dieterich (1979), Ruina (1983)

Rate- and state-dependent friction law



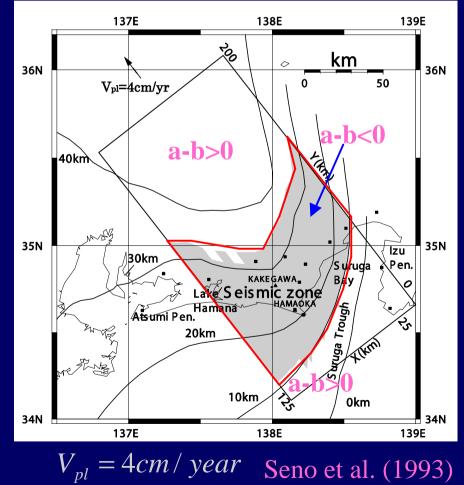
Geophysical constraints by observations

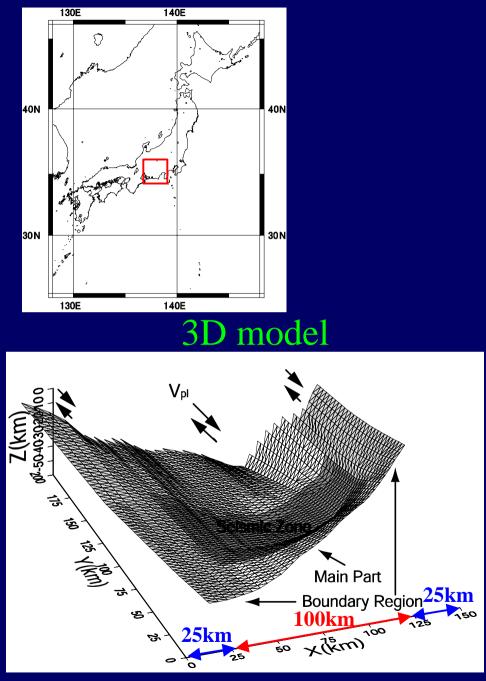
- 1. Plate configuration
 - JMA hypocentral data Harada et al. (1998)
- 2. Recurrence interval
 - 90-150 yearsIshibashi and Satake (1998)
- 3. Average seismic coupling coefficient
 - 0.5 (0-60km depth) Peterson and Seno (1984)
- 4. Crustal deformation
 - Leveling and GPSObservation by GSI
- 5. Coupling region (Locked zone or seismic zone)
 - 10-30km depth
- 6. Plate velocity
 - 4cm/year

- Matsumura (1997), Sagiya (1999) etc.
 - Seno et al. (1993)

Modeling of the plate interface

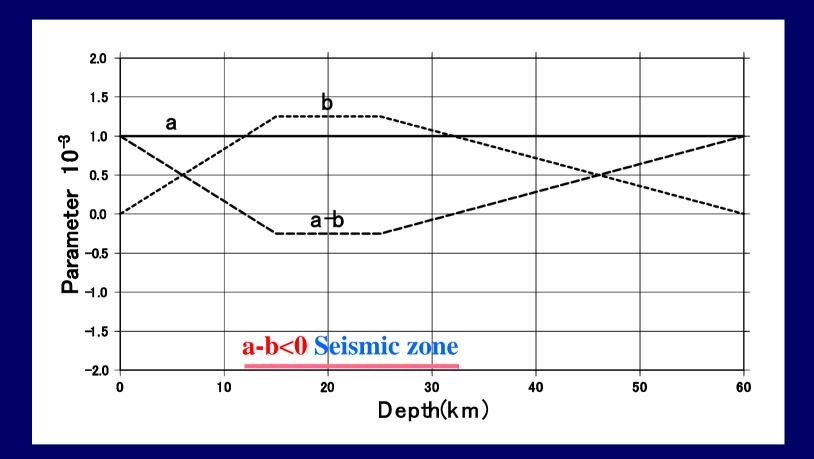
Tokai Region





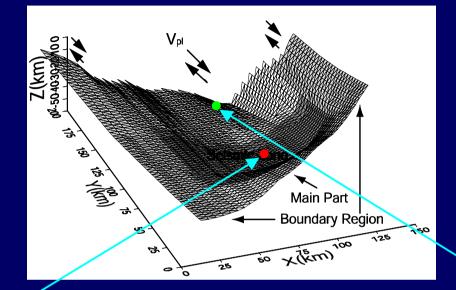
Viewed from Suruga Bay

Friction parameter



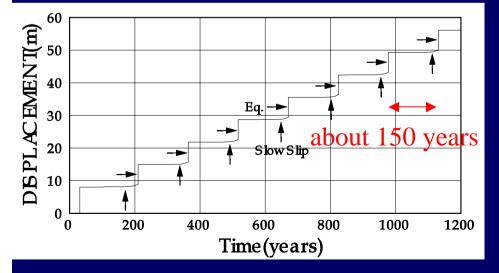
L=5cm

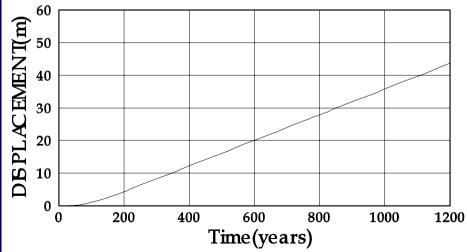
Time evolution of cumulative displacement



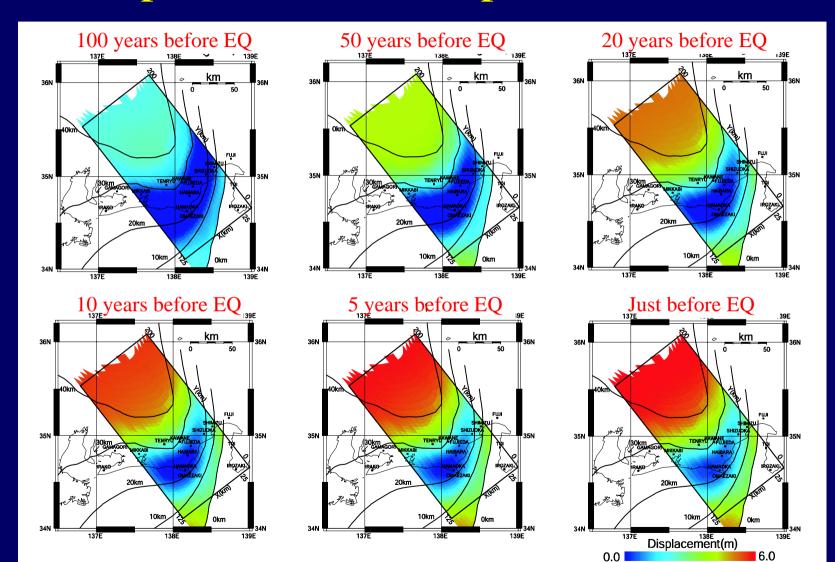
Seismic zone

Aseismic zone

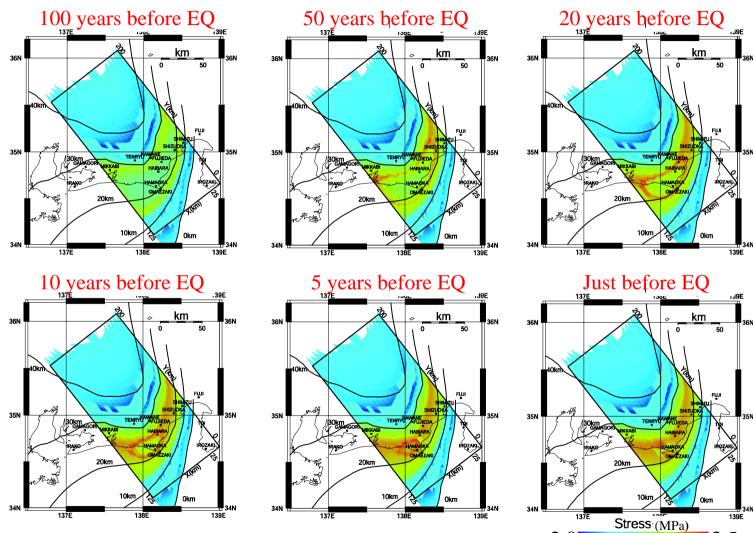




Displacement on the plate interface



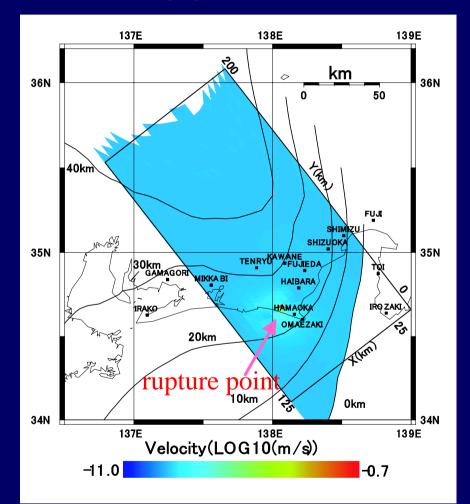
Shear stress on the plate interface



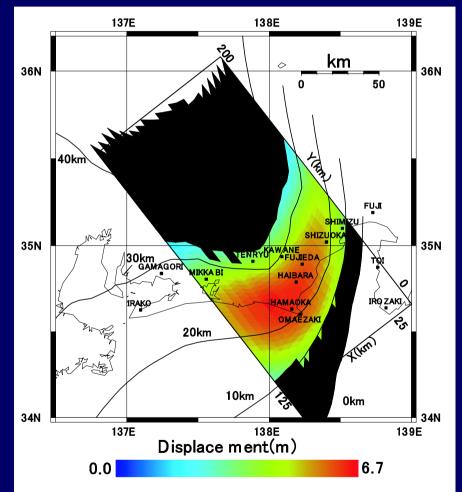
-2.0

3.5

Velocity just before EQ

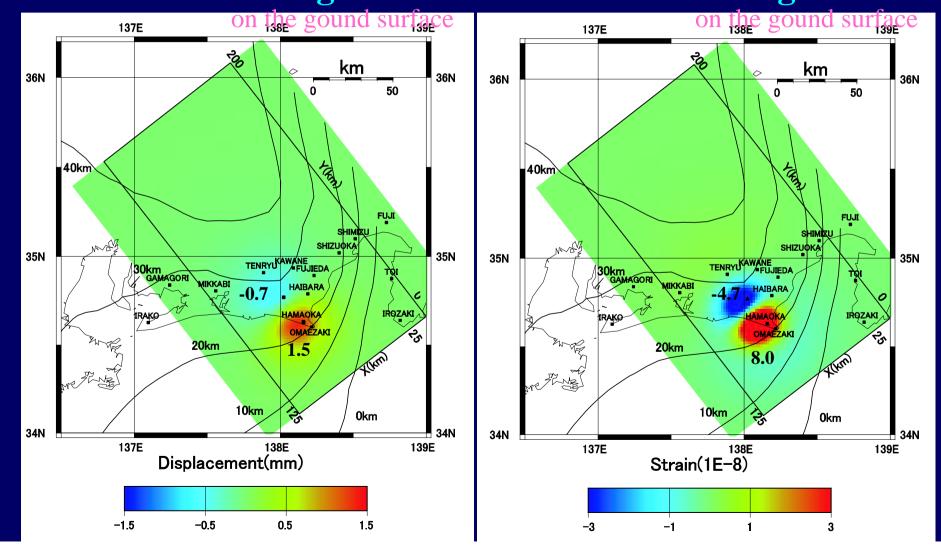


Coseismic slip distribution



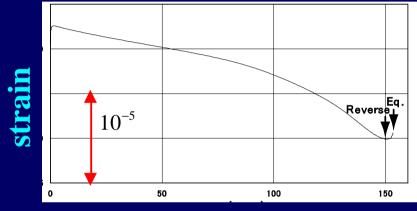
Average slip = 3.9m Mw=8.0

Level & Strain changes during one day before the earthquake Level change Strain change



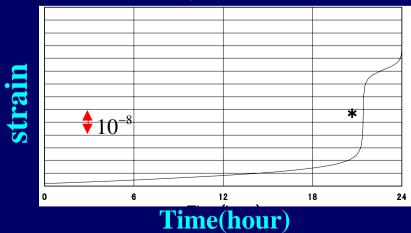
Temporal change of volumetric strain at Hamaoka station

One Cycle

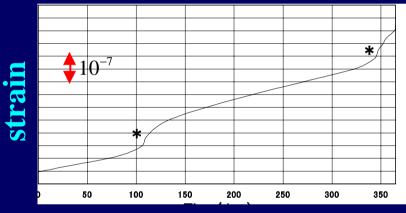


Time(year)

One day before EQ

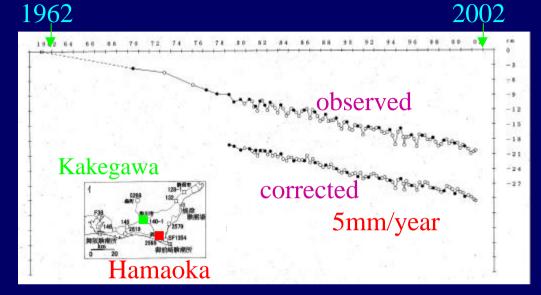


One year before EQ

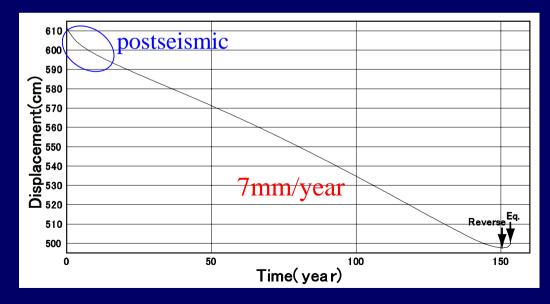


Time(day)

Subsidence at Hamaoka relative to Kakegawa

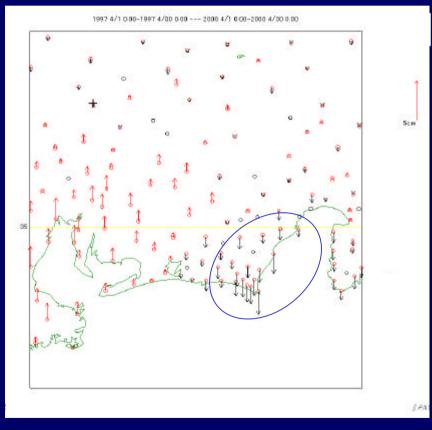


Observation by GSI

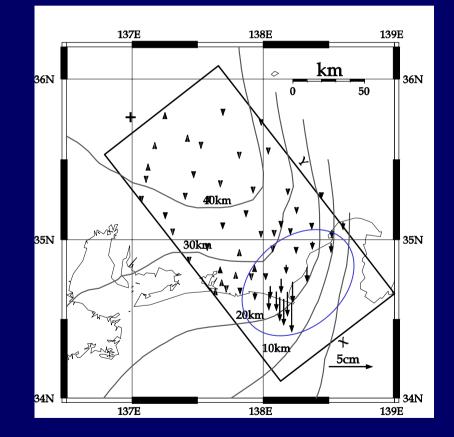


Simulation result

Comparison between GPS observation and simulation result for the vertical displacement GPS observation Simulation result

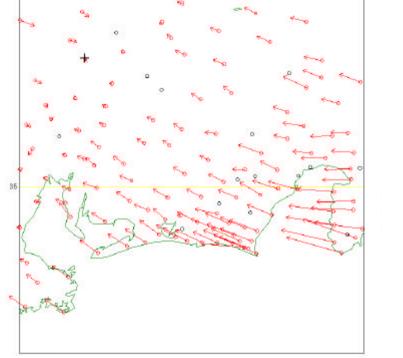


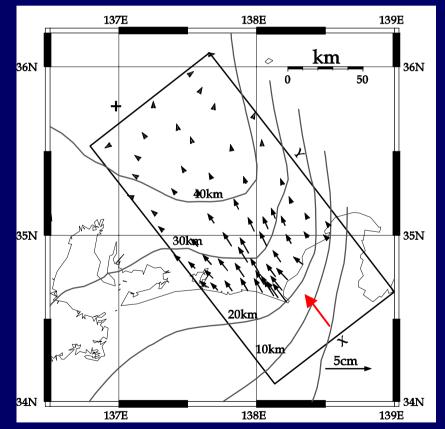




20 to 17 years before EQ

Comparison between GPS observation and simulation result for the horizontal displacement **Simulation result GPS** observation





20 to 17 years before EQ

April 1997 – April 2000

Summary

- A highly shear-stressed ring-shaped zone is formed 50 years before the earthquake around a strongly coupled region.
- Preslip gives rise to level change of the order of 1mm and volumetric strain change of 10⁻⁸ to 10⁻⁷.
- Subsidence at Hamaoka relative to Kakegawa turns to uplift several years before the earthquake.
- The results of the simulation agree well with crustal deformation observed by leveling and GPS.