Inference of regional viscosity structure from space-based geodetic data

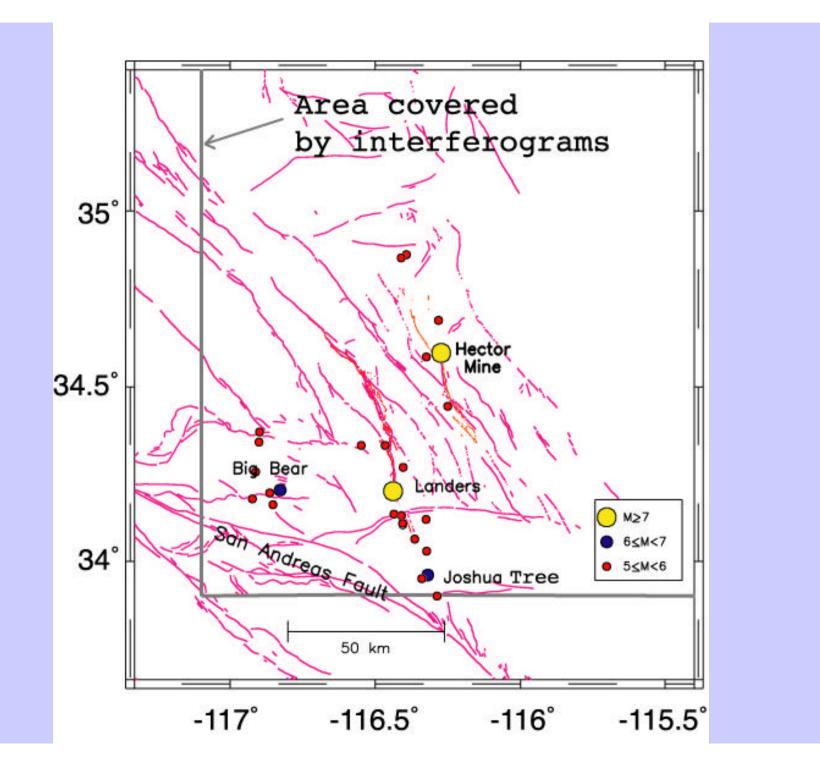
Fred F. Pollitz

U.S. Geological Survey

Menlo Park

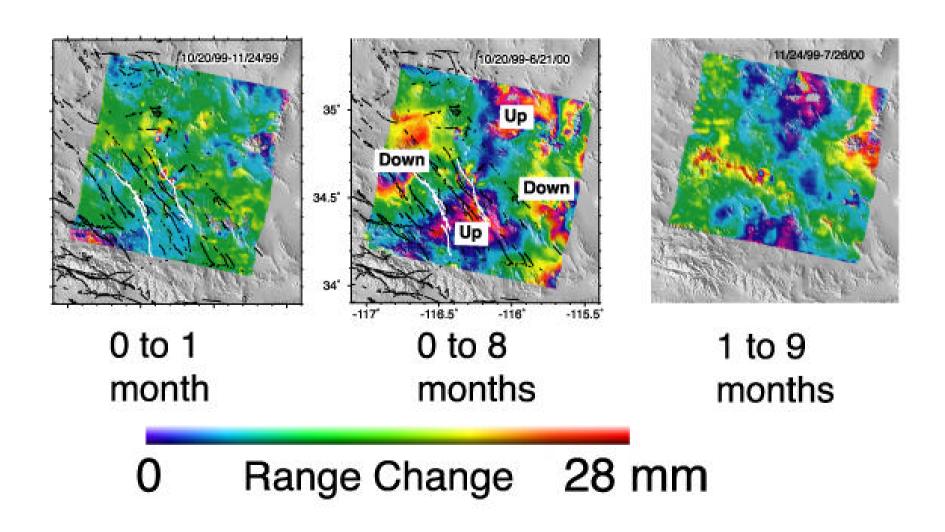
- 1. GPS and InSAR evidence for a strong crust and weak mantle beneath the Mojave Desert, southern California
- 2. Inference of transient mantle rheology beneath the Mojave Desert from three-component continuous GPS (SCIGN network) and campaign GPS (USGS network)

Acknowledgments: Jim Savage Ken Hudnut Wayne Thatcher GPS and InSAR evidence for a strong crust and weak mantle beneath the Mojave Desert, southern California

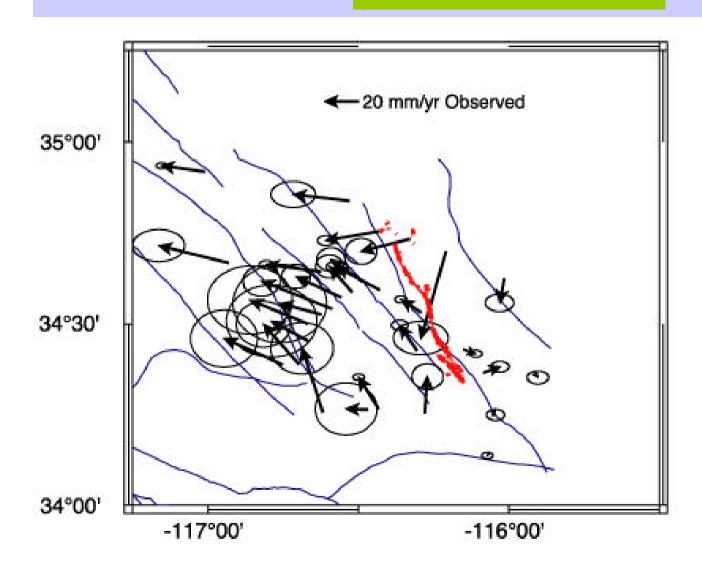


ERS interferograms provided by WINSAR

Postseismic Data

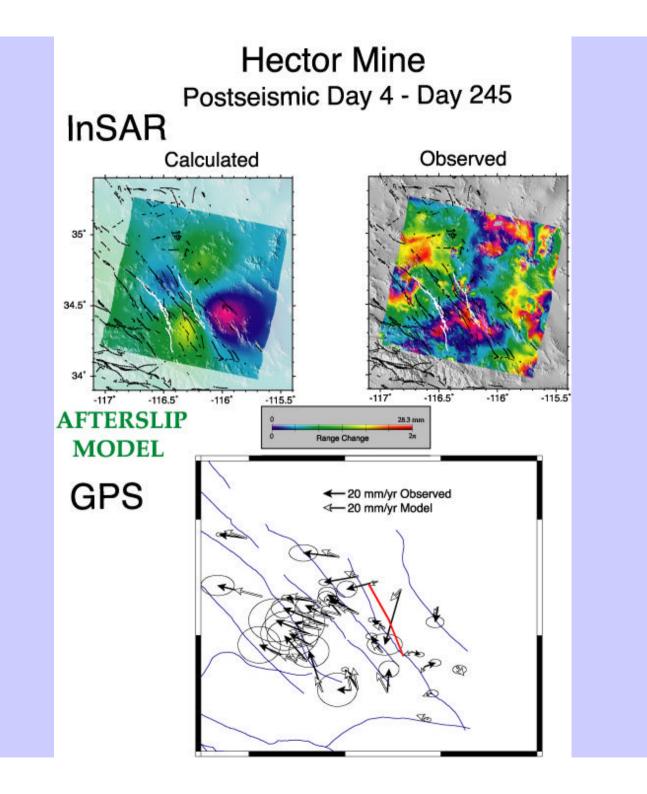


Postseismic Data

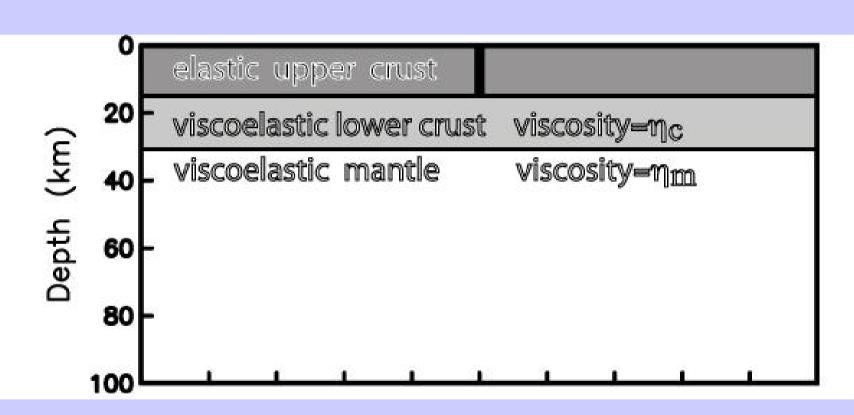


SCIGN continuous and USGS campaign GPS time series

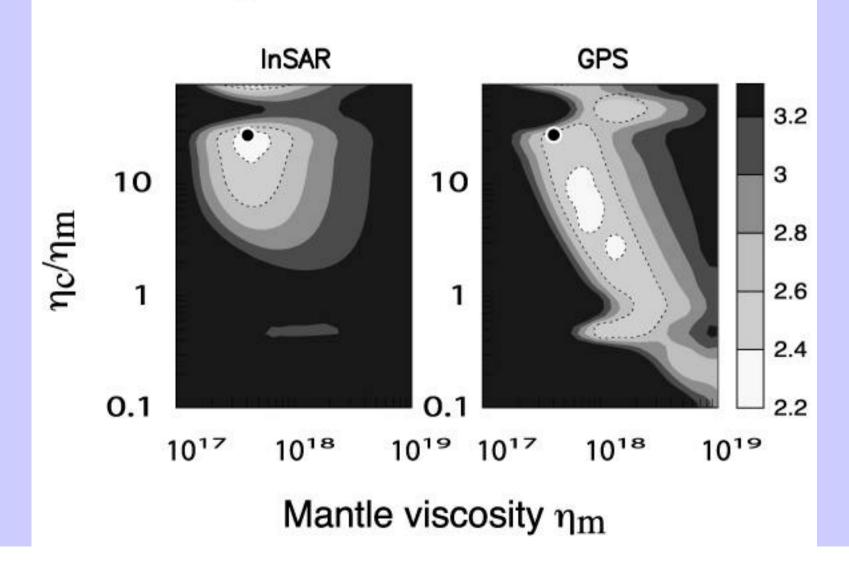
0 to 8 months



Viscoelastic Model

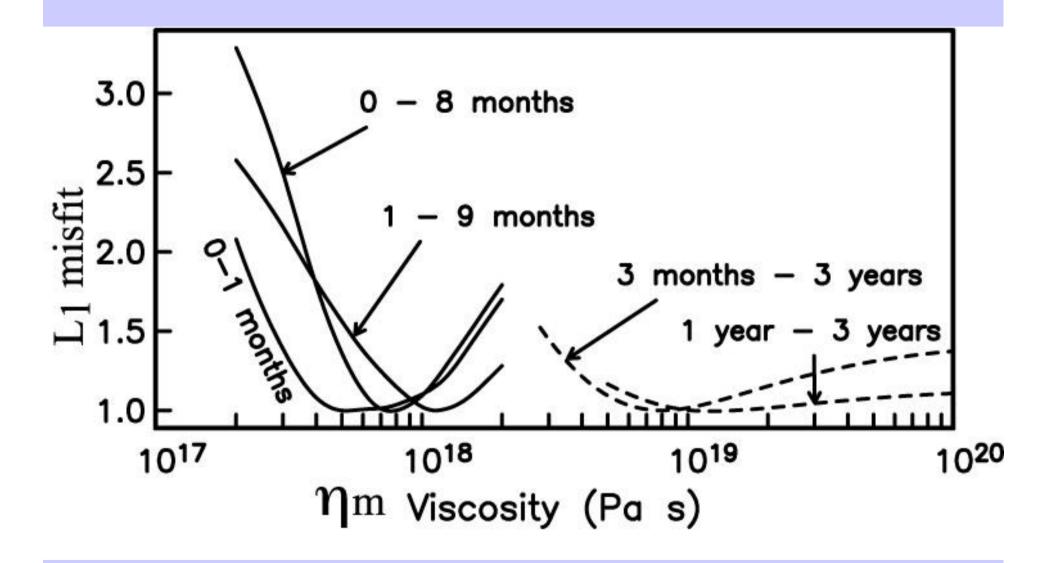


Results of grid search for η_m and η_c: log₁₀(misfit)



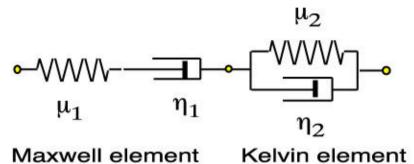
Hector Mine Postseismic Day 4 - Day 245 **InSAR** Observed Calculated 34.5 -115.5 -116.5 -116" -115.5 -116.5 VISCOELASTIC Range Change MODEL ← 20 mm/yr Observed ← 20 mm/yr Model **GPS**

Inference of transient mantle rheology beneath the Mojave Desert from three-component continuous GPS (SCIGN network) and campaign GPS (USGS network)

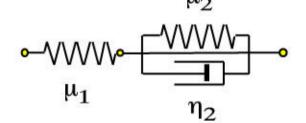


- Broadscale rapid postseismic strain after
 1992 Landers earthquake (first 3 months)
- Lower crustal afterslip (Shen et al., 1994)
- Biviscous lower crustal viscoelastic flow (Ivins, 1996)
- Broadscale elevated postseismic strain rates (several years following 1992 Landers and 1999 Hector Mine earthquakes)
- Lower crustal afterslip (Savage and Svarc, 1997; Owen et al., 2002)
- Univiscous lower crustal viscoelastic flow (Deng et al., 1998)
- Univiscous upper mantle viscoelastic flow (Pollitz et al., 2000)
- Biviscous upper mantle viscoelastic flow (Pollitz, 2002)

Burghers Body



Standard Linear Solid

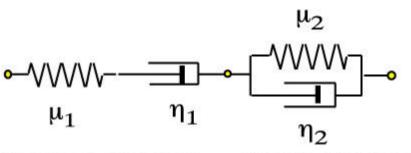


Relaxed Strength

$$\mu' = \mu_1 \, \mu_2 / (\mu_1 + \mu_2)$$

Maxwell Fluid

$$\mu_1$$
 η_1



Maxwell element

Kelvin element

Burghers Body

$$2\eta_2\ddot{\varepsilon} + 2\mu_2\dot{\varepsilon} = \frac{\eta_2}{\mu_1}\ddot{\sigma} + \left[1 + \frac{\mu_2}{\mu_1} + \frac{\eta_2}{\eta_1}\right]\dot{\sigma} + \frac{\mu_2}{\eta_1}\sigma$$

Transient Rheology

 ε = deviatoric strain

 σ = deviatoric stress

$$\dot{\varepsilon} = \frac{\sigma_0}{2\mu_1} + \frac{\sigma_0}{2\mu_2} \left[1 - \exp(-t/\tau_2) \right] + \frac{\sigma_0}{2\mu_1 \tau_1}$$

 $\dot{\epsilon}$ = instantaneous elastic response

+ transient response

+ steady state response

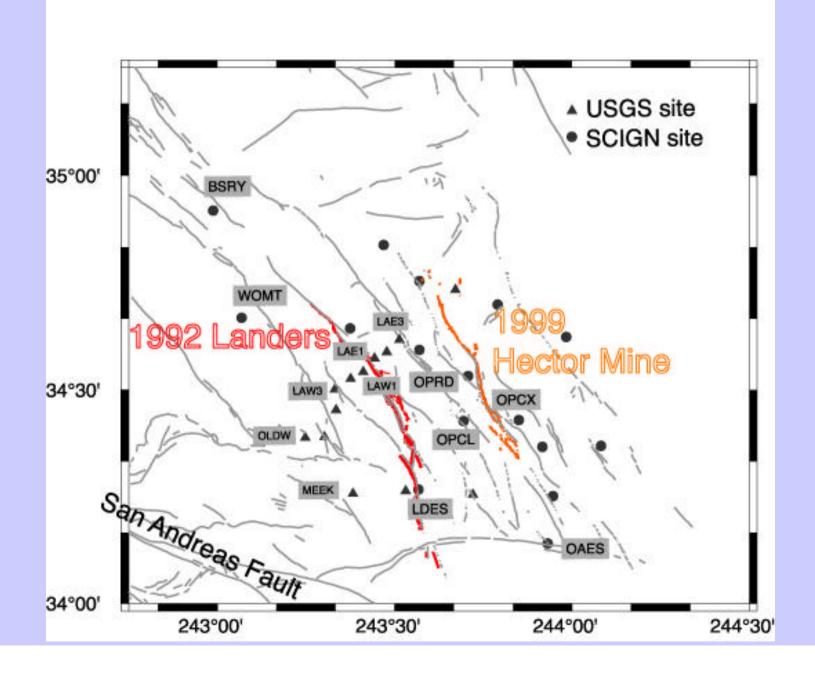
Continuous GPS Data

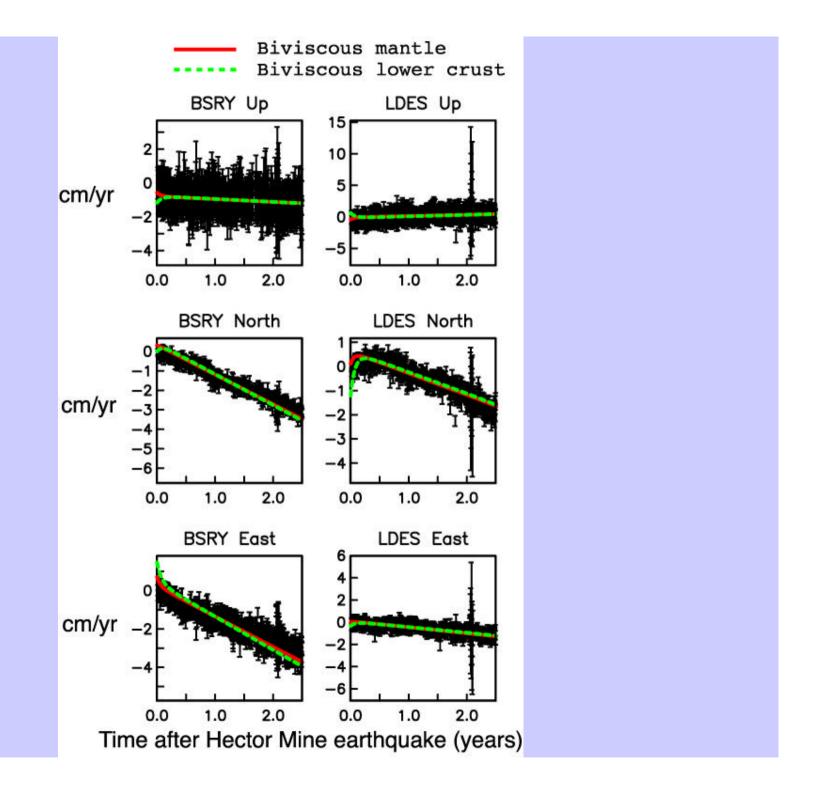
16 sites from Southern California Integrated GPS Network

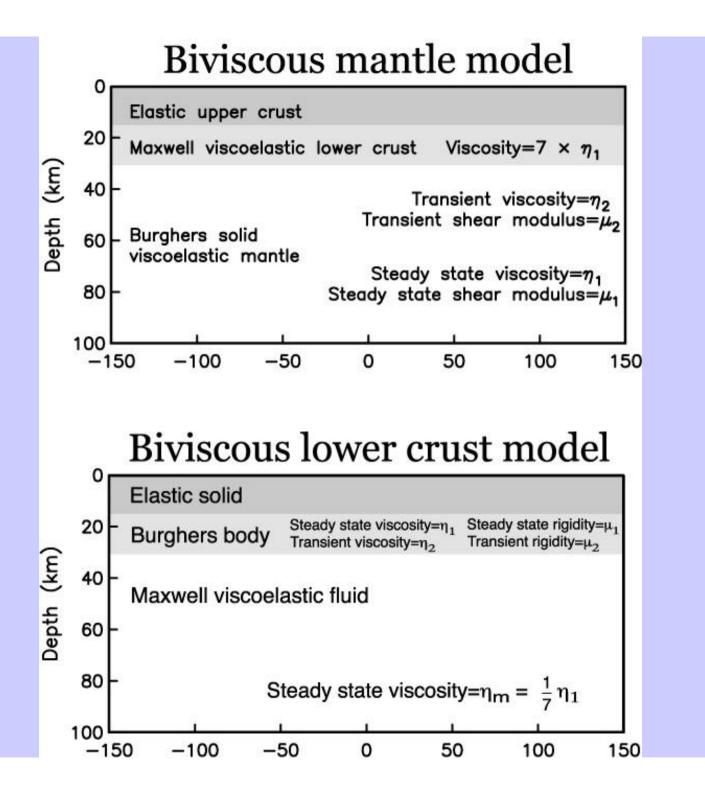
Campaign GPS Data

13 sites from U.S. Geological Survey array

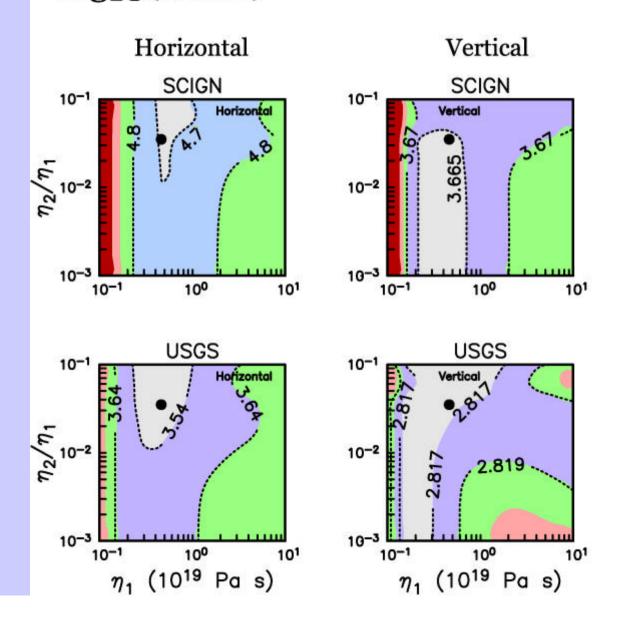
Total of 47526 samples from 87 time series (North, East, Up components)



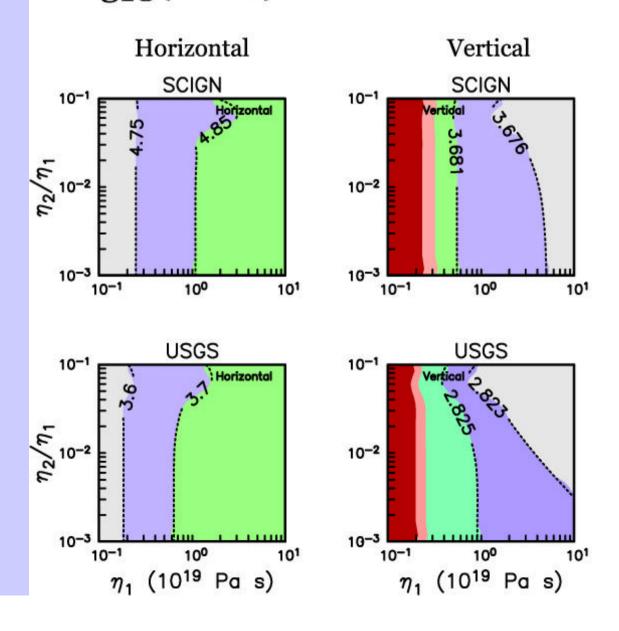




Results of grid search for η_1 and η_2 ($\mu'=0.5 \times \mu_1$): $\log_{10}(\text{misfit})$



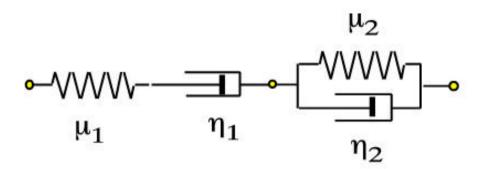
Results of grid search for η_1 and η_2 ($\mu'=0.75 \times \mu_1$): $\log_{10}(\text{misfit})$



Preferred Transient Mantle Rheology

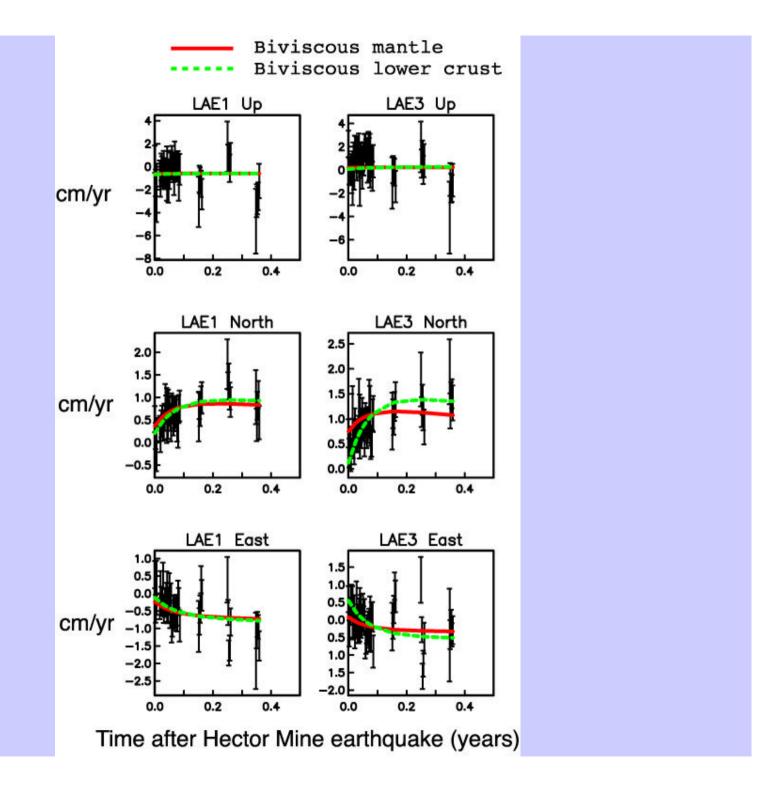
$$\eta_1 = 4 \times 10^{18} \text{ Pa s}$$

$$\eta_2 = 1.6 \times 10^{17} \text{ Pa s}$$

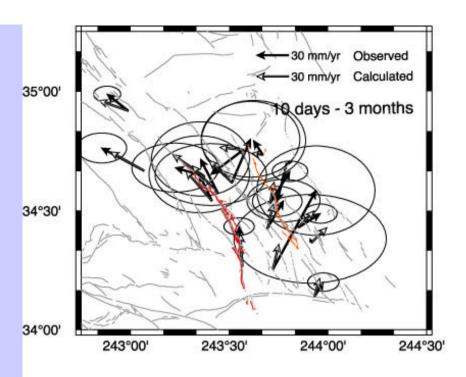


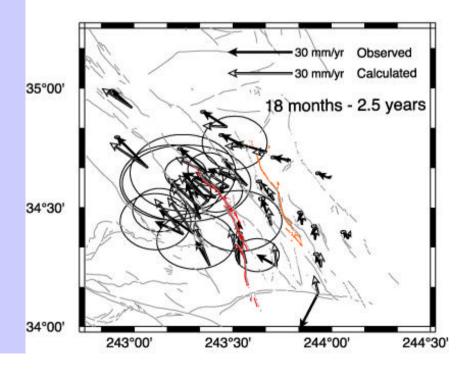
$$\mu' = \mu_1 \, \mu_2 / (\mu_1 + \mu_2) = 0.5$$

$$--> \mu_1 = \mu_2 = 70 \text{ GPa}$$

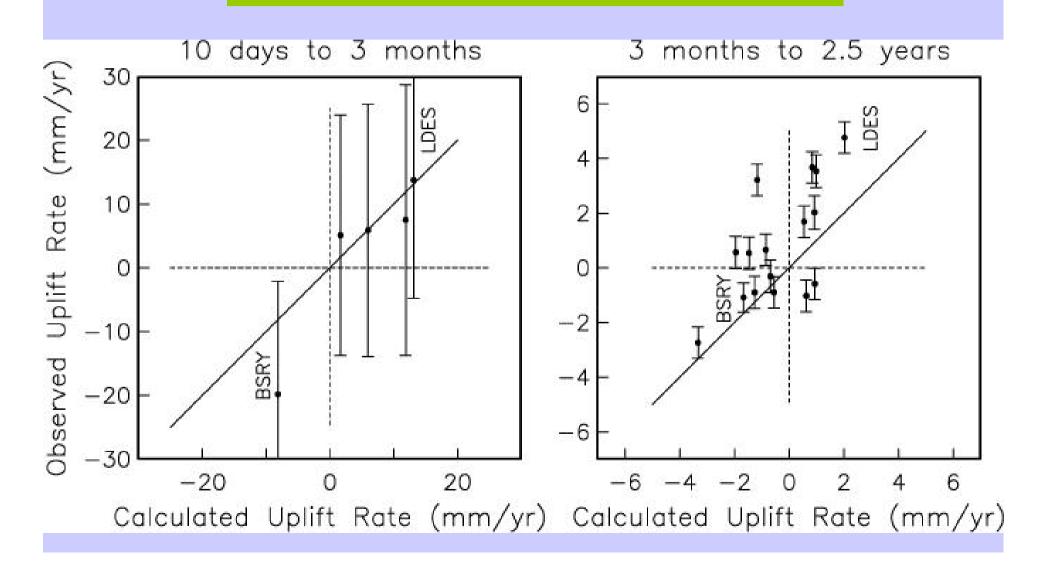


Fit of transient rheology model to horizontal velocity field



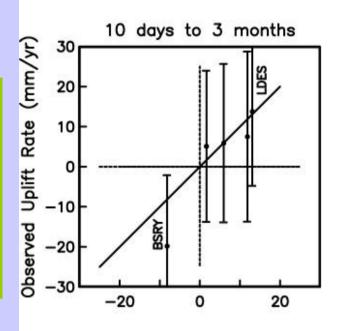


Fit of transient rheology model to vertical data

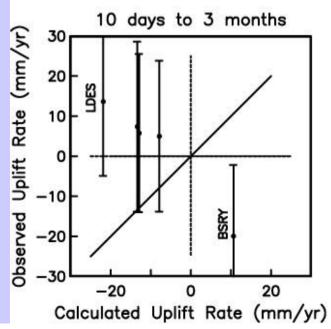


Observed & Modeled Uplift Rate

Evidence for biviscous upper mantle

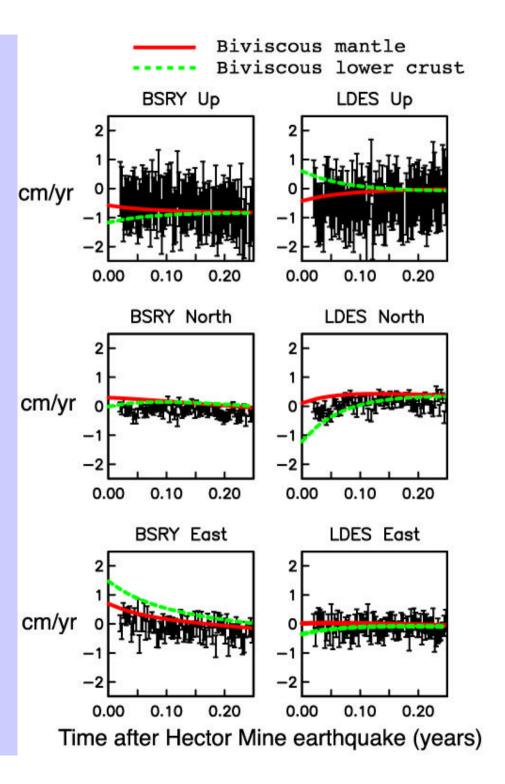


Biviscous Upper Mantle



Biviscous Lower Crust

Evidence for biviscous upper mantle



CONCLUSIONS

- InSAR data (o to 8 months) and GPS data from the first 2.5 years following the Hector Mine earthquake constrain lower crust and mantle rheology
- 2. Postseismic crustal deformation following the Hector Mine earthquake is the response of an essentially elastic crust to vigorous viscoelastic flow in the upper mantle generated by relaxation of coseismic stress changes

- The weak mantle is well characterized as a Burghers body with two material relaxation times: 0.07 years and 2 years
- 4. Constraining the details of depthdependent rheology in this study is dependent on the continuous GPS observations provided by SCIGN network